

# Natural Organ Bromine in Marine Sediments: New Evidence of Biogeochemical BR Cycling

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## Abstract

Organ bromine (Org) compounds, which are known as persistent, toxic anthropogenic pollutants, are also generated naturally in terrestrial and marine ecosystems. As a result of this research, a wide range of naturally occurring Org compounds and several enzymatic and biotic brominating routes have been uncovered. Carbon - bromine functioning, however, is unknown in the marine environment. Seabed sediments have been found to contain organic carbon (Org) and bromine (Br). Even yet, there hasn't been any clear chemical evidence that Br exists in sediments in a stable form other than the inorganic bromide (Brinorg), which is normally assumed to be conservative in marine environments. For the Org compounds found in sediments, an examination of particulate carbon from the seawater column has connected them to biologically generated Org compounds that persist as a consequence of mummification (NOM). Biogeochemical cycling of Br between organic and inorganic forms as part of the burial and degradation of NOM is evidenced by the change in Br species with sediment depth. Researchers have discovered that Br is not a conservative element in marine sediments and that the chemistry of Br and Org in marine sediments is a mystery.

**Keywords:** *natural organ, Organ bromine, Biogeochemical Br Cycling, Sediment Traps.*

## 1. INTRODUCTION

One of the primary components in saltwater is Br (bromine). Br is present in seawater at a concentration of 6.7 107 ng/L (0.84 mM), making it the fourth highest concentration known (National Astronomical Observatory of Japan 2016). Br has the second-highest seawater/crust concentration ratio after Cl (The Geochemical Society of Japan 2012). Br stays in seawater for 100 to 106 years on average (Gamo 2014). Br is categorized as a conservative element in saltwater because the vertical profiles of the Br concentration in the water column are flat (Nozaki 1992; National Astronomical Observatory of Japan 2016). However, a current investigation into the biogeochemical cycling of Br challenges the notion that Br is a conservative element (Leri et al. 2010).

Organ bromine compounds, which are known to include br, can be found in organic materials of both marine and terrestrial origins (Leri et al. 2010). More than 1600 naturally occurring organ bromine substances, many of which have marine origins, including bromoalkanes, bromopyrroles, and bromophenols, have been found (Gribble 2000). It is well known that several heterotrophic and microalgae create brominated organic chemicals (Gribble 1998). Various organisms, including bacteria, a marine annelid, and red green and brown sea algae, have been identified to have the bromoperoxidase enzymes that catalyse the synthesis of organobromine chemicals (Gribble 2000). However, nothing is known about the identities and structures of the organ bromine chemicals found in the deep sea sediments (Berg and Solomon 2016).

According to Leri et al. (2010), organic material in marine sediments is frequently brominated to the point that bromine (Br) content can be utilized to distinguish between organic material with marine and terrestrial origin (Mayer et al., 1981). Therefore, the marine cycling of this element may be significantly impacted by the removal of Br from the ocean via biological sedimentation (Channer et al., 1997). In fact, during sedimentary diagenesis, the ratios of Br to organic carbon (OC) can be fairly constant (Pedersen and Price, 1980; Mayer et al., 2007). Additionally, fluctuations in redox conditions and other parameters in sedimentary and bottomwater habitats don't seem to have much of an impact on sedimentary Br:OC ratio changes (Shimmied and Pedersen, 1990).

Bromide (Br) concentrations in seawater typically range from 60 to 80 mg kg<sup>-1</sup>, and due to brominated sedimentary organic matter, Br concentrations in marine sediments tend to be enriched in comparison (Leri et al., 2010).

## 2. LITERATURE REVIEW

**TAUANA WANKE (2015)** Acetogenins are secondary metabolites formed from the polyketide pathway, and it has long been suggested that they could serve as chemotaxonomical markers for red algae that are a part of the *Laurencia* complex. Algal C15 acetogenins differ significantly from plant C15 acetogenins in that they typically include halogens and have an enyne or a bromoallene terminal group. Laurencin and other algal acetogenins have piqued the interest of both synthetic and natural product chemists ever since they were originally discovered. This paper summarizes the research on C15 acetogenins with an emphasis on their distribution, chemical composition, and claimed biological effects.

**KARIN E. LIMBURG (2018)** the greatest artificial deoxygenated zone in the world, with rising occurrences and area-wide anoxia/hypoxia, is found in the Baltic Sea. The Baltic Sea's Atlantic cod population has declined in health, which has been primarily linked to hypoxia. In the presence of hypoxia and other reducing circumstances, Mn is accumulated in the aragonite eoliths, which are a component of the fishes' hearing and balancing system. Fish have eoliths that grow throughout their lives, therefore each fish has a lifetime record of hypoxia exposure. Otolith Mn/Ca ratios, however, are similarly susceptible to the impacts of growth. Due to the fact that Mg levels indicate metabolic activity but not hypoxia, we evaluated a new proxy to at least partially account for growth: Mn/Mg. To reconstruct lifetime histories of mean, maximum, and cumulative values of this proxy as well as others (Sr/Ca) that provide information regarding saline conditions, elemental proxies were extracted from the otoliths on a yearly basis. In order to assess fish growth and condition in connection to hypoxia experience documented by eolith proxies, we examined cod from five different time periods: the Neolithic (4500 YBP, a baseline for normoxic conditions), 1980s, 1990s, 2000s, and 2010s - under varied hypoxic intensities. Fish growth declined as their exposure to hypoxia increased, yet their condition at capture (as determined by the Fulton's K index) exhibited a strongly positive correlation with growth as evaluated by magnesium (Mg/Ca). We come to the conclusion that cod otolith chemistry proxies not only provide retrospective information on the hypoxia, growth, and metabolic health of cod throughout life, but also reflect the cod's deteriorating situation in the Baltic.

**SELVARAJ KANDASAMY (2016)** In order to completely comprehend the sources, fluxes, processes governing their cycling in various reservoirs, and their balances on a global scale, one must have a thorough understanding of the natural carbon cycle. Such a complex cycle gains another dimension as a result of anthropogenic disruptions. In order to determine whether or not these parameters are currently in balance, it is important to update the global carbon cycle by include both natural and anthropogenic sources, fluxes, and sinks throughout the land-sea continuum. Here, we compare data- and model-based estimates of terrestrial carbon fluxes to examine the export and burial rates of terrestrial organic carbon in the oceans in order to comprehend the problem of "missing terrigenous carbon." In particular for Oceania and Arctic rivers, our review reveals significant differences between field data and model output in terms of dissolved and particulate organic carbon/matter (OC/OM) fluxes and their ratios, indicating the need for further research in these areas to fine-tune the terrestrial OC export budget. We find that the marginal sediments are important sites for the burial of terrestrial OM (TOM), which is consistent with earlier opinions of Berner (1982) and Hedges and Keil. This is based on our budgeting of global sources and sinks of OC with updated estimates of marine productivity and terrestrial OM burial rate (1995). While between 60 and 80 percent of TOM is remineralized in the margins, the estimated budget also shows that the OM derived from the ocean is more efficiently remineralized than the OM derived from the terrestrial environment, highlighting the need for further advancements in the estimation of carbon burial in the marine environment. Looking back in time, increased terrestrial OC burial (by about 50%) during glacial periods compared to interglacial periods indicates a more suppressed function for continental margins and an effective transport and preservation of OM from the shelf to deep sea. Based on the review of terrestrial and marine OM burial, we recommend a number of key areas that require additional research and attention, as well as the discovery of new biogeochemical proxies and their grouping to improve future constraints on the global carbon cycle along the land-deep sea continuum.

## 3. RESEARCH METHODOLOGY

### Coastal and Deep Sea Sediments

In Barnegat Bay, NJ, and Cape Cod, MA, sediment samples (30.5 cm depth, 5.1 cm diameter; AMS soil core sampler, American Falls, ID) were collected. Removable butyrate plastic retention liners and polyethylene liner caps were used to hold the cores in place. Cores were kept horizontally at 4°C in the laboratory until further processing. The Dremel™ saw was used to cut undamaged cores into pieces 3–5 cm high for quantitative examination of Brorg and Brinorg species. Br speciation study by X-ray absorption spectroscopy was performed on portions of these sediment sections that were stored at 4°C and were still moist. The rest of the silt was air-dried for several days before being pulverised using a mortar and pestle for total Br measurement. An adhesive matrix,

NaNO<sub>3</sub>, was added to the powdered silt, and the result was a fine powder. A 13 mm die with tungsten carbide anvils under 10 tonnes of pressure in a conventional hydraulic laboratory press was used to create pellets from this combination. The butyrate plastic core liner was cut open vertically with the Dremel™ saw and pieces of the core were gently extracted and placed with little disruption of the sediment grains for X-ray spectromicroscopic examination of elemental distributions at the synchrotron beamline.

D. Sigman (Department of Geosciences, Princeton University; see Lehmann et al. [2005] for geographical and sampling details on the Bering Sea northern continental shelf "MC5" and Bower's Ridge "MC24" samples and Robinson et al. [2004] for details on the Southern Ocean "MC4" samples) kindly provided wet, presectioned, mixed sediment subsamples from Bering Sea and Southern Ocean cores. In situ X-ray absorption spectroscopy and X-ray spectromicroscopy were used to examine the Br speciation in these sediments, as well as the elemental distribution. Sediments were freeze-dried, crushed, and pelletized as previously reported [Leri et al., 2006] for total Br measurement. MC24 samples (2–3 cm and 18–19 cm) were tested before and after vigorous washing with deionizer water to determine the solubility of Br species.

#### **Sinking Particulates from Water Column Sediment Traps**

D. Sigman contributed dry, elutriated water column particles from the Sea of Okhotsk (Department of Geosciences, Princeton University). At 258 and 1061 m, sediment traps were placed every few months during a 1990–91 expedition [Honda et al., 1997]. C. Lee contributed dry, elutriated particles from the Ligurian Sea's water column (Marine Sciences Research Center, SUNY Stony Brook). Sediment traps placed at a depth of 200 metres were used to acquire this data in May 2004 [Goutx et al., 2007]. The available sediment trap material in both instances was inadequate to allow for total Br measurement, but the samples were submitted to qualitative Br speciation and elemental distribution analysis nevertheless.

#### **Quantification of Brorg and Brinorg**

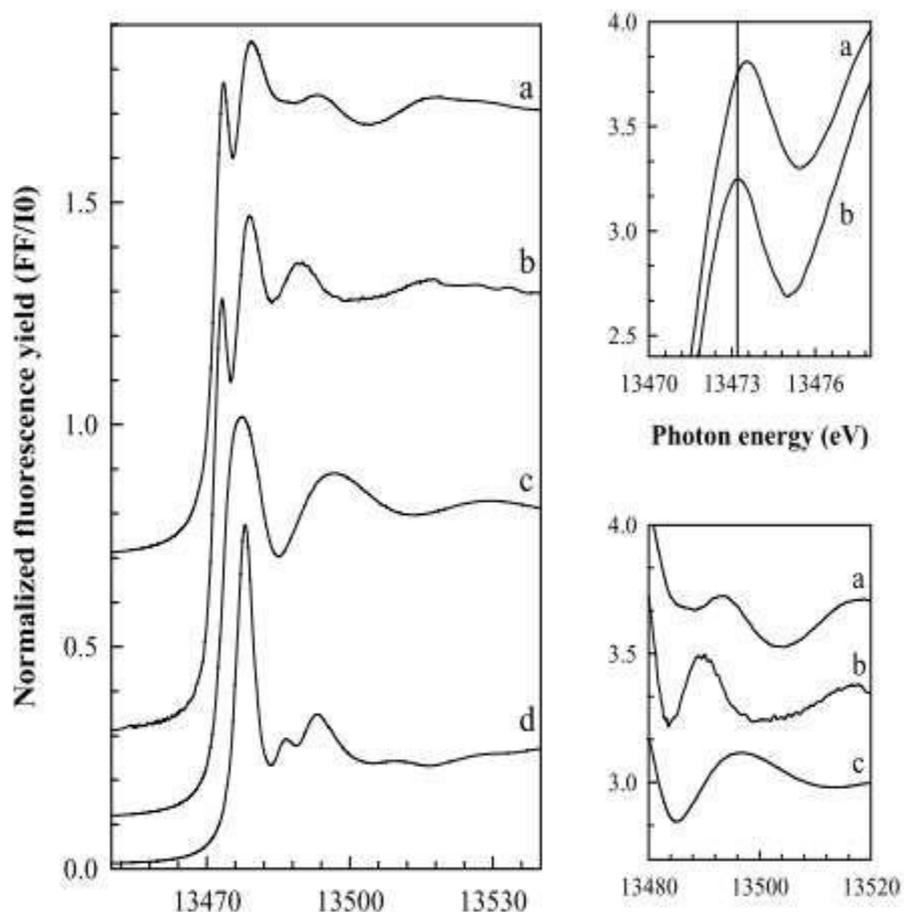
Brorg and Brinorg concentrations in sediment samples were measured using a unique quantitative approach previously disclosed [Leri et al., 2006] that combines X-ray fluorescence (XRF) and Br 1s X-ray absorption near-edge structure (XANES) spectroscopies. It was done at the National Synchrotron Light Source (NSLS, Brookhaven National Laboratory) and at beamline 43 at the Stanford Synchrotron Radiation Laboratory (SSRL, Stanford Linear Accelerator Center, and Menlo Park, CA). NSLS beam line X26A was used to obtain Br K $\alpha$  emission spectra for total Br concentration measurements.

### **4. DATA ANALYSIS**

#### **Br 1s XANES Spectroscopy**

Brinorg and Brorg species may be distinguished in natural samples without exposing them to techniques that increase the danger of unintentional bromination or partial recoveries, thanks to XANES spectroscopy.

Absorption edge characteristics in the near-edge region and the sharp rise in X-ray absorption at 13,474 eV (the Br K absorption edge) are both electronic transitions from the Br-1s shell to empty Br-4p atomic and molecular orbitals. Brinorg compounds have a greater Br transition energy as the Br oxidation state increases. Examples are Br(V) in sodium bromate (Figure 1, spectrum d) and Br(I) in water-based KBr (Figure 1, spectrum c), where the highest absorption occurs at 13,478.1 and 13,477.3 eV, respectively. In order to illustrate the relative variations in energy values, the arbitrary energy values have been used. There are distinct low-energy peaks about 13,473 eV for Brorg molecules such as 4-bromophenol (Figure 1, spectrum a) and 1-bromoeicosane (Figure 1, spectrum b). Sharp characteristics correlate to C-Br bonds and change somewhat in absorption energy depending on C-Br bond length as shown by comparisons between the aliphatic (bromoeicosane) and aromatic (bromophenol) Brorg standards (Figure 1, top right). The properties of Postedge are likewise affected by the Br coordination environment (Figure 1, bottom right). The molecular structure of brominated pollutants has been shown using theoretical modelling of X-ray absorption fine structural characteristics.



According to this method, it is possible to determine how Br<sub>inorg</sub> and Br<sub>org</sub> contribute to the overall Br content of natural samples by fitting the spectra of typical model compounds to those of natural samples. It is possible to learn about the local coordination environment of Br via Br 1s XANES spectroscopy; however, this technique does not reveal the overall size and structure of Br-containing molecules in natural samples.

#### Geochemistry of Br as a Function of Sediment Depth and C<sub>org</sub> Content

In coastal and marine sediment cores, chemical profiles show how Br<sub>org</sub> concentration changes as a function of sediment depth and NOM content (Table 1). Speciation shifts from Br<sub>org</sub>-dominated to Br<sub>inorg</sub>-dominated in opal-rich Southern Ocean Antarctic strata formed of Holocene diatomaceous ooze [Robinson et al., 2004]. The Br 1s XANES transition narrows and amplifies as a result of this change, and the Br<sub>org</sub> shoulder disappears from the spectrum. Aqueous Br<sub>inorg</sub> is similar to the KBr (aq) standard. Below the sediment-water contact, the absolute Br<sub>org</sub> concentration drops precipitously, and is now 14–16 cm below the detection limit. In contrast to Br<sub>org</sub>, the concentration of C<sub>org</sub> decreases more gradually increasing depth. There is a correlation between depth and the Br<sub>org</sub>:C<sub>org</sub> ratio (Table 1), which indicates that the Br<sub>org</sub> in these sediments is rather labile.

**Table 1 Br Speciation, Concentration, and Total Organic Carbon in Sediments**

Depth Below Sediment-Water Interface (cm)	Br <sub>inorg</sub> (ppm)	Br <sub>org</sub> (ppm)	Total [Br] (ppm)	Total [C <sub>org</sub> ] (g/kg)	$\frac{\text{Br}_{\text{org}}}{\text{C}_{\text{org}}}$ Ratio/ $10^{-3}$
	Antarctic Pacific				
0–0.5	201	108	309	ND	ND
0.5–1	108	36	144	5.5 ± 0.1	6.5

2–2.5	126	14	140	$4.3 \pm 0.2$	3.3
7–8	190	10	200	$3.3 \pm 0.2$	3.0
14–16	130	<1	130	$2.6 \pm 0.1$	UD
Bering Sea (Northern Continental Shelf)					
2	91	59	150	$13.1 \pm 0.1$	4.5
7	79	60	139	$12.0 \pm 0.3$	5.0
13	66	28	94	$11.7 \pm 0.3$	2.4
19	83	28	111	$10.8 \pm 0.1$	2.6
Bering Sea (Bower's Ridge)					
2–3	162	8	170	ND	ND
4–5	118	6	124	$7.2 \pm 0.1$	0.8
8–9	52	42	94	$8.9 \pm 0.1$	4.7
18–19	29	16	45	$8.7 \pm 0.9$	1.8
28–29	27	33	60	$18.0 \pm 0.8$	1.8
34–35	71	47	118	$10.9 \pm 0.4$	4.3
Cape Cod, MA, Estuary					
0–3.5	12	18	30	$4.3 \pm 0.8$	4.2
4–9	7	3	10	$0.9 \pm 0.2$	3.3
9–16	8	3	11	$1.1 \pm 0.3$	2.7
16–20	7	3	10	$0.4 \pm 0.1$	7.5
20–25	8	<1	8	$0.3 \pm 0.1$	UD
Barnegat Bay, NJ					
7–9 cm (organic lens included)	4	2	6	$0.7 \pm 0.1$	2.9

Biological productivity of the Bering Sea continental shelf, combined with the steep slope's physical properties that route significant amounts of NOM into the sediments below [Lehmann et al., 2005]. It is possible that the higher Brorg concentrations in these sediments than in Antarctica's at comparable depths may be explained by the fact that the Corg concentration drops progressively with increasing sediment depth but is still high at >10 cm (Table 1). Once again, it seems that the Brorg component in the NOM is a comparably labile one in the Bering Sea northern continental shelf location.

Rising total Br concentrations in coastal Louisiana sediments were utilized as a proxy for increasing marine (as opposed to terrestrially produced) NOM content and our data corroborate those findings. As sediment depth

drops, the total Br:Corg ratio declines, which the authors interpreted as an increase in the decomposition of marine NOM during diagenesis in Louisiana strata.

## 5. CONCLUSION

The organic component of estuarine and marine sediments from diverse geological settings is dominated by stable Brorg, according to these findings. Covalent C-Br bonds are likely responsible for the observed Br-Corg interactions in marine sediments. According to recent studies of terrestrial peat bogs, Br integrated into geomacromolecules such as humics (e.g., Brorg) provides for a large reservoir of this halogen inside the pedosphere (Biester et al., 2006, 2004). There is an abundance of Brorg in environmental samples, which suggests that it is involved in a complex biogeochemical cycle involving the formation of C-Br bonds, humification, and the eventual regeneration of Brinorg through the degradation of NOM, as well as in the chemistry of specific microenvironments. To reevaluate the designation of Br as a conservative element in seawater, the scale of Br cycling must be examined in the context of the global marine Br budget. Further research is needed to determine the relative importance of biotic and abiotic bromination processes. This research might have a significant impact on how we evaluate the effects of human-made substances having chemical characteristics similar to those found in natural Brorg.

## 6. REFERENCE

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